

MODELING OF VERTICAL PROPAGATION OF ATMOSPHERIC ACOUSTIC DISTURBANCE INITIATED BY A PULSE AT THE LOWER BOUNDARY

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Initial equations

System of hydrothermodynamics

$$\frac{\partial \vec{V}}{\partial t} = -\frac{\vec{\nabla} p'}{\bar{\rho}} + \vec{g} \frac{\rho'}{\bar{\rho}}, \quad (1)$$

$$\frac{\partial p'}{\partial t} = -\vec{V} \cdot (\vec{\nabla} \bar{p}) - \gamma \bar{p} (\vec{\nabla} \cdot \vec{V}), \quad (2)$$

$$\frac{\partial \rho'}{\partial t} = -\vec{V} \cdot (\vec{\nabla} \bar{\rho}) - \bar{\rho} (\vec{\nabla} \cdot \vec{V}), \quad (3)$$

where \vec{V} is the gas flow velocity; p' is the wave contribution to the background pressure \bar{p} ; ρ' is the wave contribution to the background density $\bar{\rho}$.

$$\bar{\rho}(z) = \rho_0 \exp\left(-\frac{z}{H}\right), \quad \bar{p}(z) = p_0 \exp\left(-\frac{z}{H}\right) = \bar{\rho}(z)gH, \quad (4)$$

where $\rho_0 = \bar{\rho}(z = 0)$, $p_0 = \bar{p}(z = 0)$, and the relationship between the equilibrium pressure and density follows from the stationary zero-order equality $d\bar{p}/dz = -g\bar{\rho}(z)$, given $g_x = 0$, $g_y = 0$, $g_z = -g$.

Initial equations

1D system of hydrothermodynamics

The one-dimensional linearized system of hydrothermodynamics:

$$\frac{\partial U}{\partial t} = \frac{1}{\rho_0} \left(\frac{\gamma - 2}{2\gamma H(0)} - \frac{H(z)}{H(0)} \frac{\partial}{\partial z} \right) P + \frac{\Phi}{\gamma H(0) \rho_0}, \quad (5)$$

$$\frac{\partial P}{\partial t} = -\gamma g H(0) \rho_0 \frac{\partial U}{\partial z} - g H(0) \rho_0 \frac{\gamma - 2}{2H(z)} U, \quad (6)$$

$$\frac{\partial \Phi}{\partial t} = -\frac{\gamma - 1 + \gamma \frac{dH(z)}{dz}}{H(z)} g H(0) \rho_0 U, \quad (7)$$

Here P, Φ, U are functions that represent the pressure perturbation p' , the entropy perturbation $\varphi' = p' - \gamma \rho' \bar{p} / \bar{\rho}$ and the vertical flow velocity V respectively and are related to the real values as

$$P = p' \cdot \exp \left(\int_0^z \frac{dz'}{2H(z')} \right), \quad \Phi = \varphi' \cdot \exp \left(\int_0^z \frac{dz'}{2H(z')} \right), \quad U = V \cdot \exp \left(- \int_0^z \frac{dz'}{2H(z')} \right),$$

$\gamma = C_p / C_v$; C_p, C_v are molar heat capacities at constant pressure and volume, $g = g_z$ is the vertical component of gravity field vector \vec{g} , ρ_0 and ρ' is the air density at the lower boundary its perturbation respectively.

Initial equations

1D Klein-Gordon equation

The one-dimensional system of hydrothermodynamics (1)-(3) can be reduced to the one-dimensional Klein-Gordon equation by differentiating the equation (5) with respect to the time t and replacing the derivatives $\partial P/\partial t$ and $\partial \Phi/\partial t$ with (6) and (7):

$$\frac{\partial^2 U}{\partial t^2} - \gamma g H(z) \frac{\partial^2 U}{\partial z^2} + \frac{g\gamma}{4H(z)} \left(1 + 2 \frac{dH(z)}{dz} \right) U = 0 \quad (8)$$

or

$$\frac{\partial^2 U}{\partial t^2} - c^2(z) \frac{\partial^2 U}{\partial z^2} + a(z) U = 0, \quad (9)$$

where

$$c(z) = \sqrt{\gamma g H(z)}, \quad a(z) = \frac{g\gamma}{4H(z)} \left(1 + 2 \frac{dH(z)}{dz} \right). \quad (10)$$

Initial equations

Initial and boundary condition

We complete Klein-Gordon equation with the initial-boundary conditions:

$$U \Big|_{t=0} = U_t \Big|_{t=0} = 0, \quad U \Big|_{z=0} = F^n(t), \quad (11)$$

$$F^n(t) = \frac{A_n}{(n+1)!} \lambda^{n+2} t^{n+1} e^{-\lambda t}, \quad \text{at } t > 0, \quad F(t) = 0, \quad \text{at } t \leq 0, \quad (12)$$

where λ characterizes the duration of the pulse and A its amplitude.

Later we will discuss results obtained with the boundary conditions for $n = 0, 1, 2$:

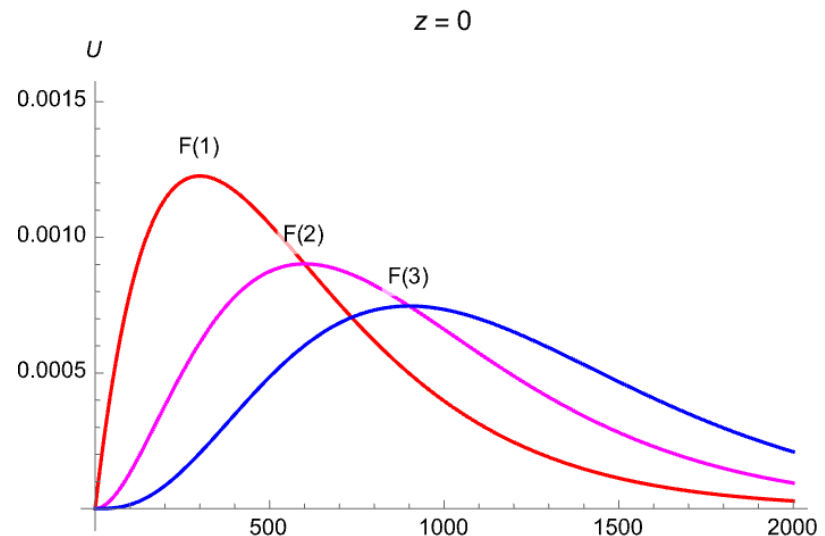


Figure: Boundary conditions $F^1(t) = A\lambda^2 t e^{-\lambda t}$ (red), $F^2(\tau) = \frac{A}{2} \lambda^3 t^2 e^{-\lambda t}$ (pink), $F^3(\tau) = \frac{A}{6} \lambda^4 t^3 e^{-\lambda t}$ (blue) for $A = 1, \lambda = 1/300$.

Solution

Analytical solution in case $H = H_0$

Theorem 1. For constant value of the atmospheric scale height $H(z) = H(z = 0) = H_0$ the coefficients take on a constant value $c(0)$, $a(0)$ and the initial-boundary value problem for the Klein-Gordon equation can be analytically solved [1]:

$$U(t, z) = U_1(t, z) + U_2(t, z) \quad (13)$$

$$U_1(t, z) = \frac{1}{\pi} \operatorname{Re} \int_0^{\sqrt{a(0)}} \mathcal{F}(i\eta) e^{i\left(\eta t - \frac{z}{c(0)} \sqrt{a(0) - \eta^2}\right)} d\eta, \quad (14)$$

$$U_2(t, z) = \frac{1}{\pi} \operatorname{Re} \int_{\sqrt{a(0)}}^{\infty} \mathcal{F}(i\eta) e^{i\left(\eta t - \frac{z}{c(0)} \sqrt{\eta^2 - a(0)}\right)} d\eta, \quad (15)$$

where $\mathcal{F}(i\eta)$ is the Laplace image of the boundary condition (11):

$$\mathcal{F}(i\eta) = \mathcal{F}(s) = \int_0^{\infty} F(t) e^{-st} dt. \quad (16)$$

Remark 1: For $a = 0$, the resulting formula (13) becomes the exact formula of the initial-boundary value problem solution for the wave equation.

[1] Smirnova, E.S. *Asymptotics of the Solution of an Initial–Boundary Value Problem for the One-Dimensional Klein–Gordon Equation on the Half-Line*. Math Notes 114, 608–618 (2023). <https://doi.org/10.1134/S0001434623090286>

Solution

$H=H(z)$ approximation

Since finding a solution in the case of $H = H(z)$ requires its explicit form, for the altitude range from 0 to 200 km, the following approximation was chosen:

$$H(z) = 7000 + 0.135z \tanh^5\left(\frac{z}{12000}\right). \quad (17)$$

Remark 2: Such an approximation at low altitudes gives an almost constant value of both the atmospheric scale height $H = H(z)$ and of the main coefficients $c(z)$ and $a(z)$ of the Klein-Gordon equation. Thus, in this area the solution of the initial boundary value problem in the case of variable coefficients $c(z)$ and $a(z)$ coincides with the exact solution for constant values of the coefficients $c(0)$ and $a(0)$.

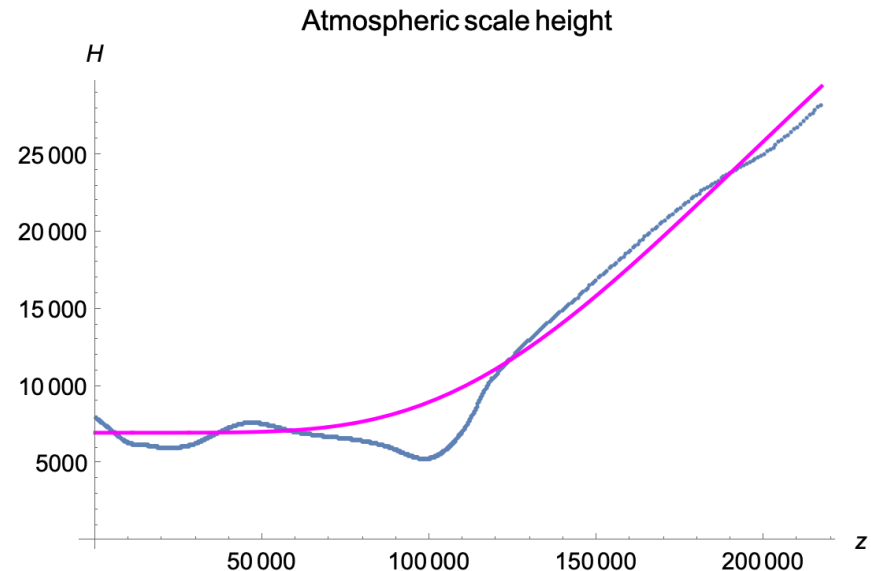


Figure: Atmospheric scale height $H = H(T(z))$ calculated from numerical simulation data for temperature (blue) and its approximation (17) (pink).

Solution

Asymptotic solution in case $H = H(z)$

Theorem 2: The leading term of the formal asymptotic solution of the problem (9),(11) is determined by the formulas [2]:

$$U = U_1 + U_2, \quad U_1(\tau, y) = \frac{e^{-\frac{y}{h}}}{\pi\tau} \operatorname{Re} [\mathcal{F}(0, h)], \quad U_2 = \frac{c(y)}{c(0)} (K_{A_\tau^+} A_0^+ + K_{A_\tau^-} A_0^-). \quad (18)$$

Remark 3: For low altitudes in the region under study, the form of the asymptotic of the wave-propagating part of the solution:

$$\begin{aligned} U_2(y, \tau) &= K_{A_\tau^+} A_0^+ + K_{A_\tau^-} A_0^- = \frac{1}{2\pi h} \int_0^\infty \bar{\mathcal{F}}(\sqrt{c^2(0)p^2 + a(0)}, h) e^{\frac{i}{h}(py - \tau\sqrt{c^2(0)p^2 + a(0)})} \frac{c^2(0)p}{\sqrt{c^2(0)p^2 + a(0)}} dp - \\ &\quad - \frac{1}{2\pi h} \int_{-\infty}^0 \mathcal{F}(\sqrt{c^2(0)p^2 + a(0)}, h) e^{\frac{i}{h}(py + \tau\sqrt{c^2(0)p^2 + a(0)})} \frac{c^2(0)p}{\sqrt{c^2(0)p^2 + a(0)}} dp = \\ &= \frac{1}{\pi h} \operatorname{Re} \int_0^\infty \mathcal{F}(\sqrt{c^2(0)p^2 + a(0)}, h) e^{\frac{i}{h}(\tau\sqrt{c^2(0)p^2 + a(0)} - py)} \frac{c^2(0)p}{\sqrt{c^2(0)p^2 + a(0)}} dp. \end{aligned} \quad (19)$$

For medium and high altitudes in the region under study, the form of the asymptotics of the wave-propagating part of the solution:

$$\begin{aligned} U_2 &= \frac{c(y)}{c(0)} (K_{A_\tau^+} A_0^+ + K_{A_\tau^-} A_0^-) = \frac{1}{\sqrt{2\pi h}} \frac{c(y)}{c(0)} \left(\frac{e^{\frac{i}{h}S^+(\alpha, \tau)} \alpha c^2(0)}{\sqrt{J^+(\alpha, \tau)} (\alpha^2 c^2(0) + a(0))} \bar{\mathcal{F}}(\sqrt{\alpha^2 c^2(0) + a(0)}, h) \right) \Big|_{\alpha=\alpha^+(y, \tau)} + \\ &\quad + \frac{1}{\sqrt{2\pi h}} \frac{c(y)}{c(0)} \left(\frac{e^{-\frac{i}{h}S^-(\alpha, \tau)} \alpha c^2(0)}{\sqrt{J^-(\alpha, \tau)} (\alpha^2 c^2(0) + a(0))} \mathcal{F}(\sqrt{\alpha^2 c^2(0) + a(0)}, h) \right) \Big|_{\alpha=\alpha^-(y, \tau)} = \\ &= \sqrt{\frac{2}{\pi h}} \frac{c(y)}{c(0)} \operatorname{Re} \left(\frac{e^{\frac{i}{h}S^+(\alpha, \tau)} \alpha c^2(0)}{\sqrt{J^+(\alpha, \tau)} (\alpha^2 c^2(0) + a(0))} \bar{\mathcal{F}}(\sqrt{\alpha^2 c^2(0) + a(0)}, h) \right) \Big|_{\alpha=\alpha^+(y, \tau)}. \end{aligned} \quad (20)$$

[2] Dobrokhotov, S., Smirnova, E. *Asymptotics of the Solution of the Initial Boundary Value Problem for the One-Dimensional Klein–Gordon Equation with Variable Coefficients*. Russ. J. Math. Phys. 31, 187–198 (2024).

Results

Asymptotic solution in case $H = H_0$

Return to the physical functions of flow velocity disturbance is carried out through (8):

$$V = U \cdot \exp\left(\frac{z}{2H_0}\right) \quad (21)$$

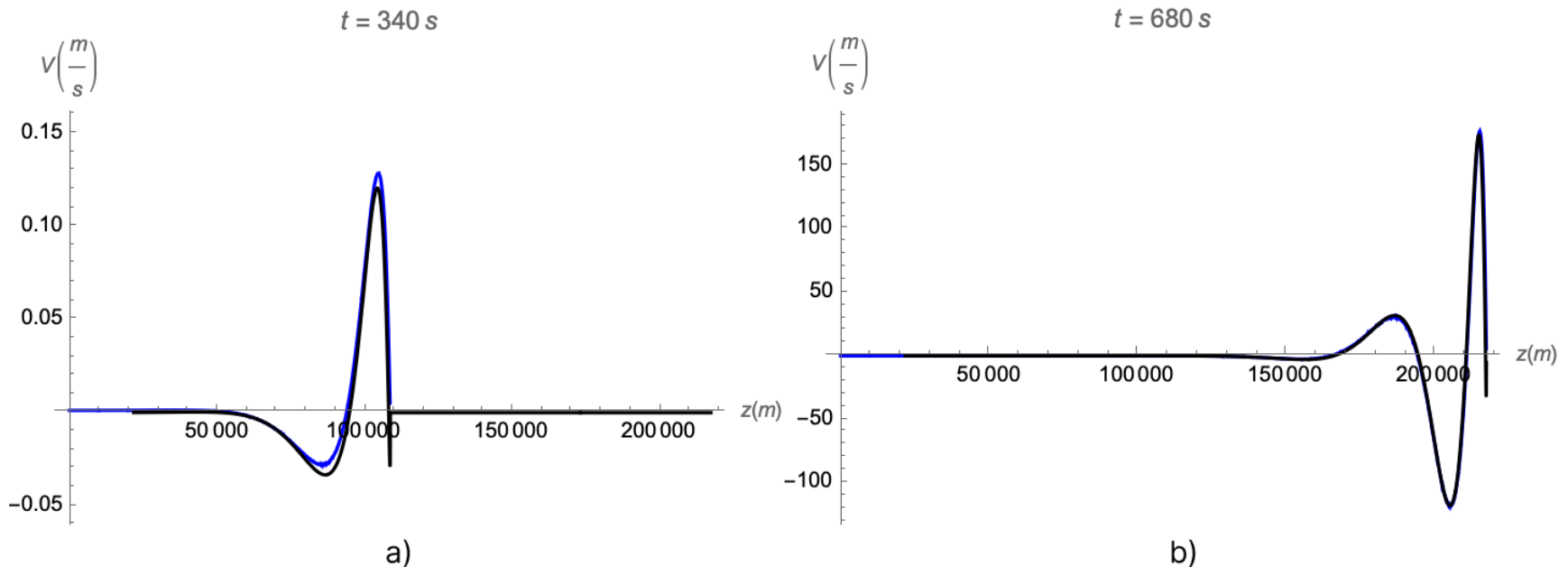


Figure: Comparison of analytical (numerical calculation) and asymptotic solutions for constant value $H = H_0 = 7000 \text{ m}$.

Results

Asymptotic solution in case $H = H(z)$

Return to the physical functions of flow velocity disturbance is carried out through (8):

$$V = U \cdot \exp \left(\int_0^z \frac{dz'}{2H(z')} \right) \quad (22)$$

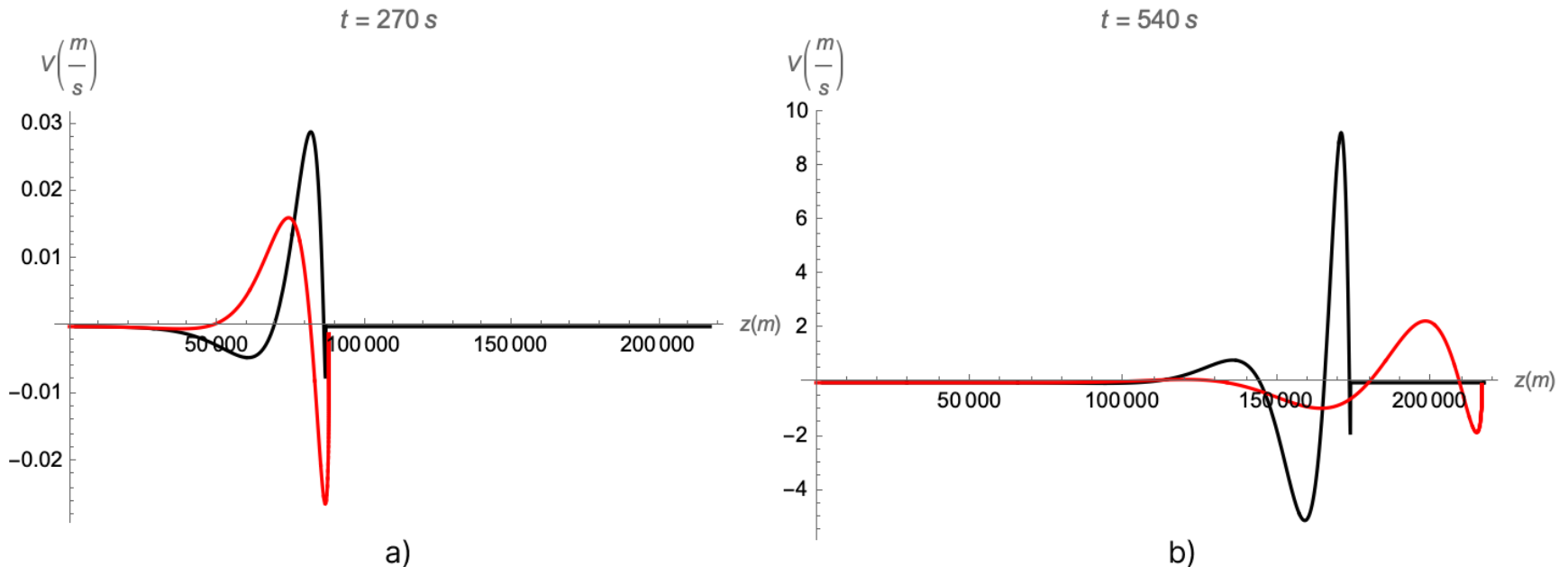


Figure: Comparison of time evolution of the asymptotic solutions (19) for $H = H_0 = 7000 m$ (black) and (20) for $H = H(z)$ (17) (red).

Results

Different boundary conditions

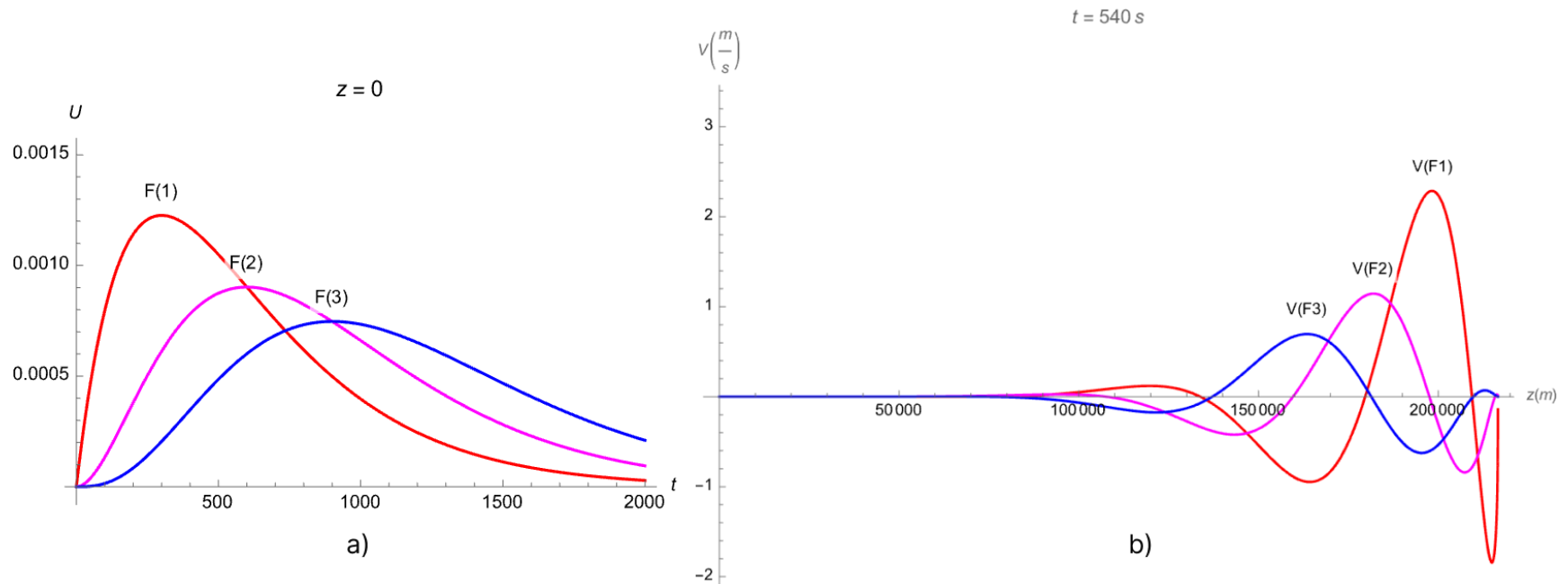


Figure: **(a)** Boundary conditions $F^1(t)$ (red), $F^2(t)$ (pink), $F^3(t)$ (blue); **(b)** Asymptotics of solution at $t = 540$ s in case $H = H(z)$ corresponding each boundary condition. To simplify the analysis, the amplitudes were normalized: $A_1 = 1$; $A_2 = 10$; $A_3 = 100$.

Discussion

The initial-boundary value problem is solved both analytically and asymptotically in a general form, therefore, parameters of the problem, as well as the boundary condition, can be refined for a more specific physical problem.

Thank you for your attention.

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