

The Influence of Aerosol's  
Size Dispersion  
on Dissipative Instability of  
Aerosol Flow in  
Planetary Atmospheres' Plasma

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# Dusty (aerosol or complex) plasma:

neutrals, electrons  
and ions

+

large charged particles of  
micron/submicron size  
(dust/aerosols).

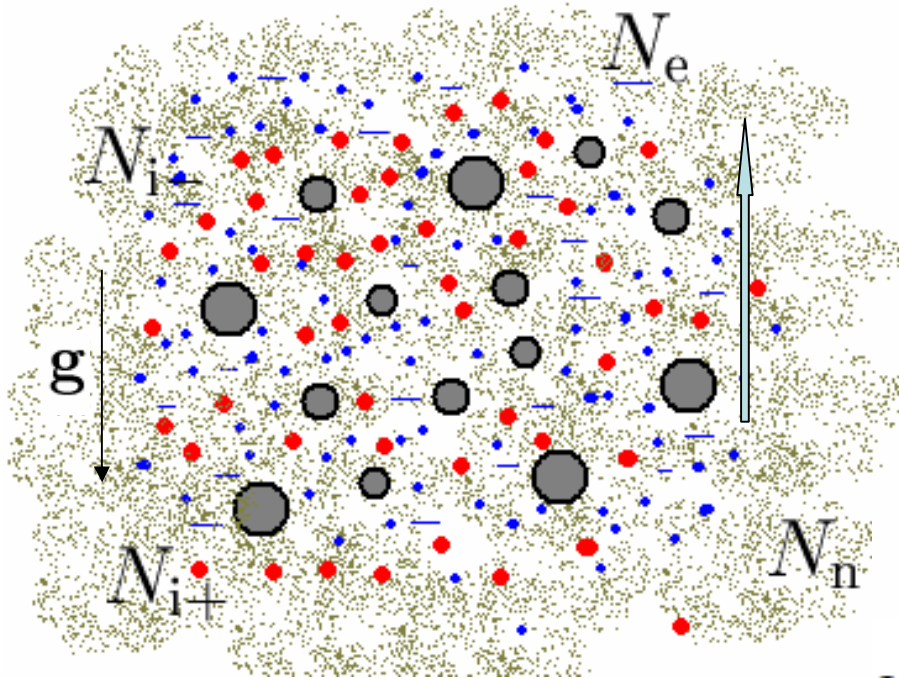
Charging of the large particles  
is an additional non-stationary  
process.

Dusty plasma can be found in  
space, different regions of  
planet atmospheres, is studied  
in laboratories.

The presence of dust causes phenomena with new spatial and time scales, including **dust-acoustic wave** (DAW). The development of DAW instability can generate fine structures in electron density/space charge/electric field in dusty plasma.

In the Earth's atmosphere such fine structures can be found in thunderclouds ( $\sim 10$  m), or in mesopause region ( $\sim 5$  cm-5 m, causing Polar Mesospheric Summer Echoes).

# Aerosol (dust) particles in weakly-ionized collisional plasma:




Low temperatures  $T \sim 300$  K;  
thermal dispersion of velocities is negligible.

Aerosol flow is caused by gravity,  
velocity is determined by friction force:

$$V = \frac{g}{\nu}; \quad \nu = \frac{2\sqrt{2\pi}}{\pi} \gamma_F \frac{1}{\rho_a R} \frac{N_n T_n}{V_{T_n}} = \nu^* \frac{a}{R};$$

$$V = V^* \frac{R}{a} \quad \nu^* \equiv \nu(R = a),$$

$$V^* \equiv V(R = a)$$

:  $Q_d(R) \mathbf{E}; \quad M(R) \mathbf{g}; \quad \mathbf{F}_{fr}$

$Q_d(R)$  и  $M(R)$  – charge and mass of a particle with radii  $R$ .

$\nu$  – effective frequency of aerosol collisions with neutral gas

# Aerosol's size dispersion (model distributions):

Monodispersed ensemble:

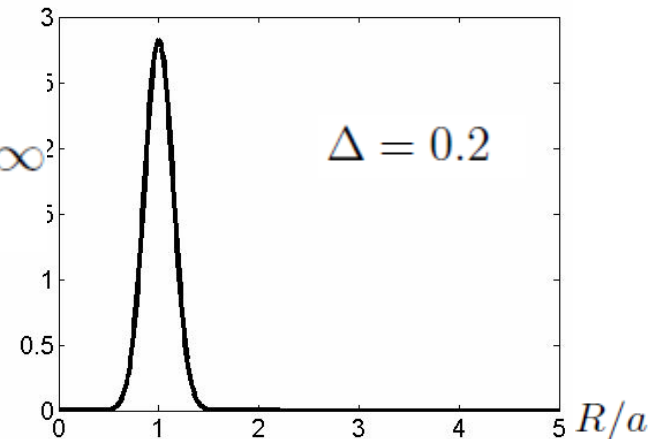
$$f_R = \delta(R - a), \quad 0 < R < \infty$$

Gaussian distribution:

$$f_R = \frac{a}{\delta a \sqrt{\pi}} \exp \left[ - \left( \frac{R - a}{\delta a} \right)^2 \right], \quad 0 < R < \infty$$

Radii should be positive and finite:

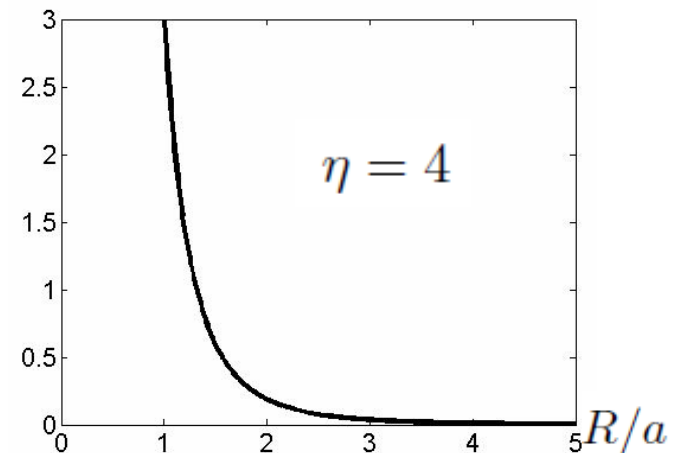
$$\Delta = \frac{\delta a}{a} < 0.5$$



Power distribution:

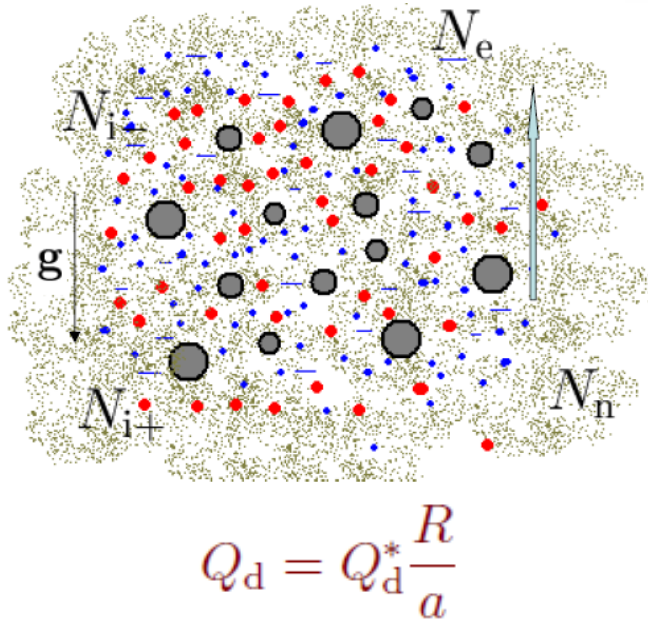
$$f_R = \frac{\eta - 1}{a} \left( \frac{R}{a} \right)^{-\eta}, \quad a < R < \infty.$$

Flow integrals converge:  $\eta > 3$



## Charging of Aerosols

(due to collisions with ions and electrons, limited orbit model):



$$\frac{dQ_d}{dt} = eR^2 \sqrt{8\pi} \left[ n_i V_{T_i} \left( 1 - \frac{eQ_d}{RT_i} \right) - n_e V_{T_e} \exp \frac{eQ_d}{RT_e} \right]$$

Charging by high-energy electrons:

$$n_h \ll n_e < n_i, T_h \gg T_e \quad n_h \ll n_i \quad I_h > I_e$$

$$\frac{dQ_d}{dt} = eR^2 \sqrt{8\pi} n_i V_{T_i} \left( 1 - \frac{eQ_d}{RT_i} \right) - I_h$$

**Ions and electrons:** hydrodynamic, collisions with neutrals+electric field.

Relative motion of charged components (aerosols and ions) in collisional plasma can cause a dissipative instability of Dust Acoustic Wave.

# Dispersion equation:

(time evolution of quasi-static electric disturbances)

For processes  $\sim \exp(-i\omega t + i\mathbf{k}\mathbf{r})$ ,  $E_0=0$

$$1 - 4\pi N_d^0 \int \frac{Q_d^2 f_R dR}{M \Delta_d} + \frac{4\pi\sigma_i}{\Delta_i} \left[ 1 + iN_d^0 \int \frac{G_i^0 f_R dR}{\omega - \mathbf{k}\mathbf{V}_{an} + i\nu_{ch}} \right] +$$

$$+ \frac{4\pi\sigma_e}{\Delta_e} \left[ 1 + i\delta N_d^0 \int \frac{G_i^0 f_R dR}{\omega - \mathbf{k}\mathbf{V}_{an} + i\nu_{ch}} \right] + 4\pi e N_d^0 n_+^0 \left( \frac{1}{\Delta_i} - \frac{1}{\Delta_e} \right) \int \frac{G_i^0 Q_d f_R dR}{M \Delta_d} = 0$$

here

$$\Delta_d = (\omega - \mathbf{k}\mathbf{V}_{an})(\omega - \mathbf{k}\mathbf{V}_{an} + i\nu)$$

$$\Delta_i = \omega + iD_i k^2 + i\nu_R, \quad \Delta_e = \omega + iD_e k^2 + i\delta\nu_R \quad \delta = n_+^0/n_-^0$$

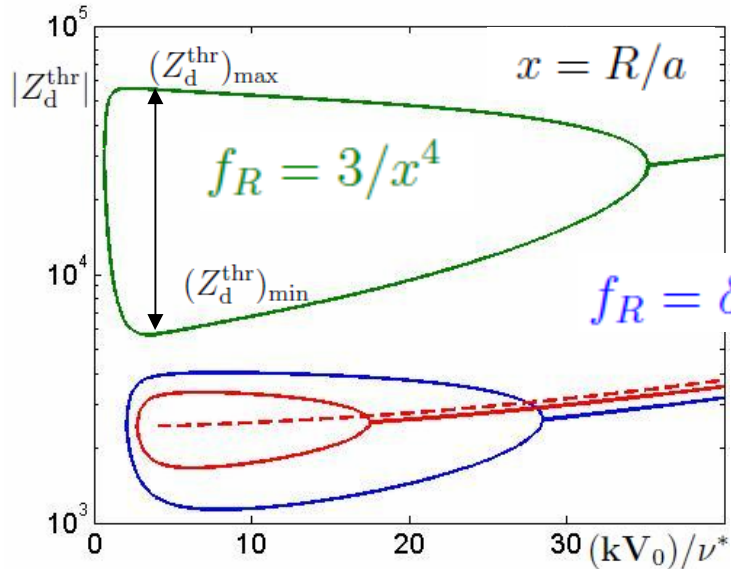
$\nu = g/V_{an}$  - effective collision frequency

$D_i, D_e$  - diffusion coefficients  $\sigma_i, \sigma_e$  - conductivities

$\nu_{ch} = \frac{d}{dQ_d} (G_e n_e - G_i n_i) |_{Q_d=Q_D^0}$  - reverse charging time;

$G_i n_i, G_e n_e$  - ion/electron flows  $\nu_R$  - ion/electron collisions with dust

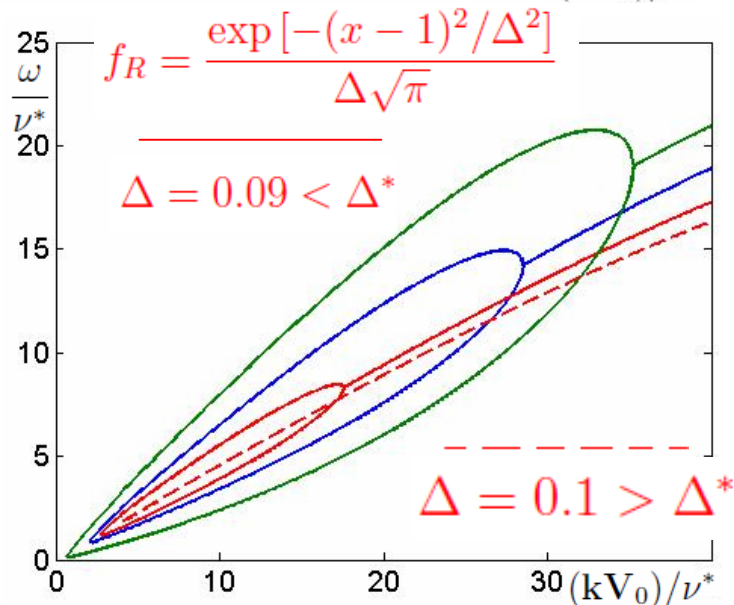
# Instability Threshold (solution of the dispersion equation if $\text{Im}\omega=0$ ):



Dispersion equation has two solutions on the threshold, so the sufficient condition for the instability is:

$$(Z_d^{\text{thr}})_{\text{min}} < Z_d < (Z_d^{\text{thr}})_{\text{max}}$$

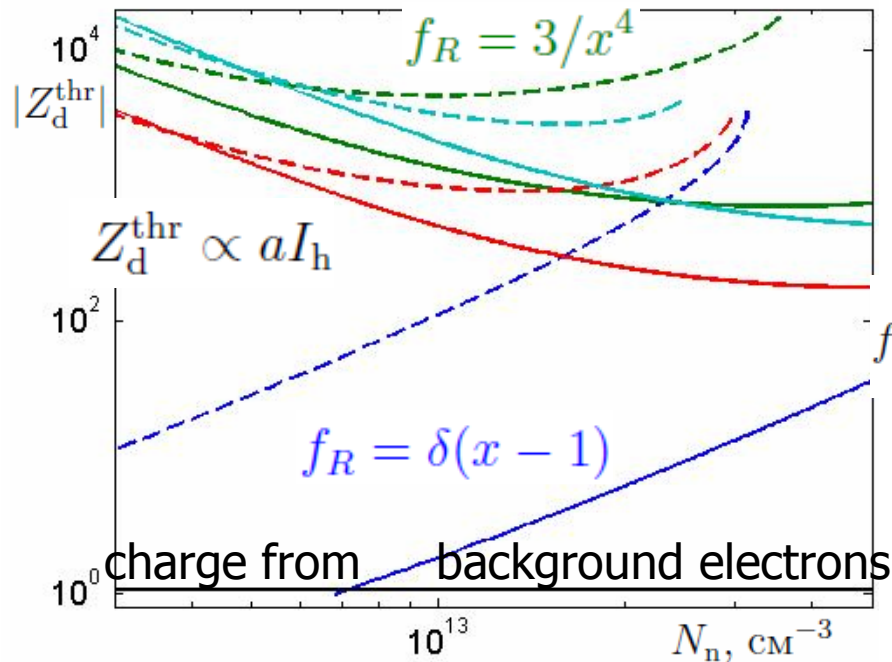
Power distribution increases interval between  $(Z_d^{\text{thr}})_{\text{min}}$  и  $(Z_d^{\text{thr}})_{\text{max}}$ ; Gaussian distribution decreases it. If  $\Delta \geq \Delta^*$ , instability is absent.



$$Z_d e = Q_d^*$$

Threshold value minimum over wavenumbers defines optimal scales.

# Earth's mesosphere (heights 80-90 km):



Neutrals:  $N_2$ ,  $T \sim 120$  K;

Ions: heavy clusters  $H(H_2O)_n$

Aerosols: ice,  $a = 100$  nm

$g = 9.8$  m/c<sup>2</sup>

$\Delta = 0.09$

$$f_R = \frac{\exp[-(x-1)^2/\Delta^2]}{\Delta\sqrt{\pi}}$$

$\Delta = 0.2$

$\rho_a = 1$  g/cm<sup>3</sup>,  $\delta = n_i/n_e = 10$ ,

$\nu_i/N_n = 9 \cdot 10^{-10}$  cm<sup>3</sup>/c

$\rho_a = 2$  g/cm<sup>3</sup>,  $\delta = n_i/n_e = 50$ ,

$\nu_i/N_n = 27 \cdot 10^{-10}$  cm<sup>3</sup>/c

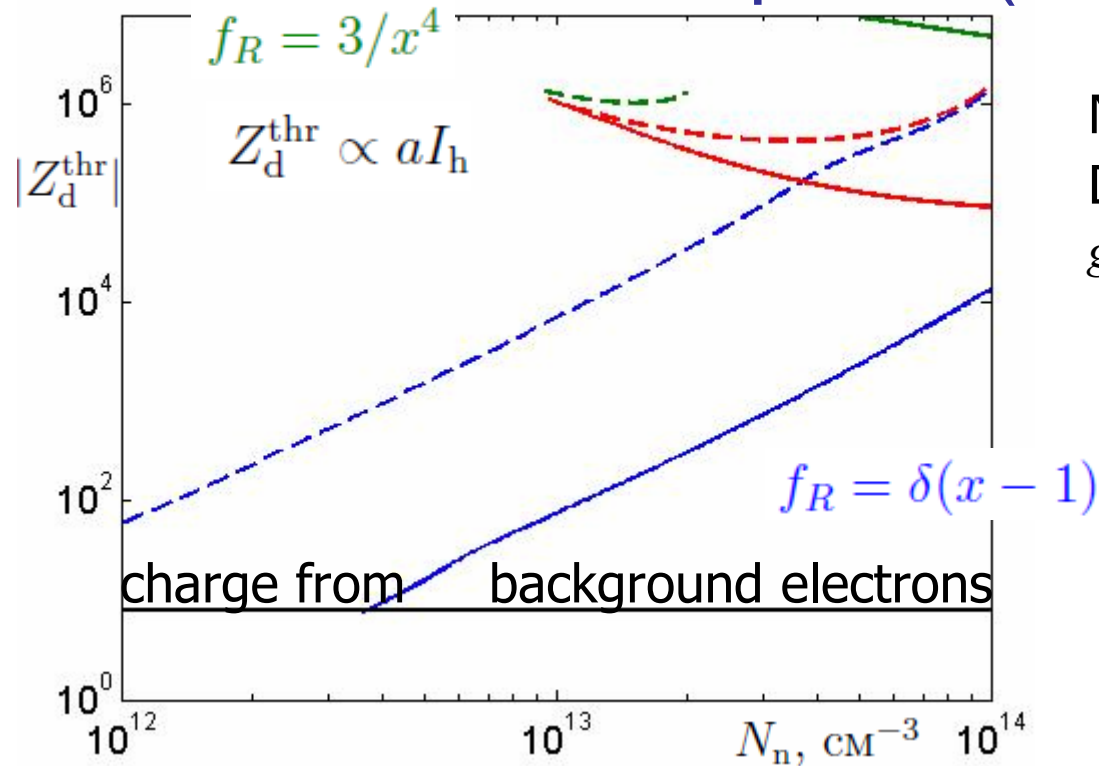
$\lambda_{opt} = 20 \div 45$  cm     $\lambda_{opt} = 60 \div 100$  cm     $\lambda_{opt} = 50 \div 150$  cm     $\lambda_{opt} = 80 \div 200$  cm

Size dispersion **increases** the instability threshold and optimal scales significantly and also changes threshold's dependence on some of the parameters. The influence of size dispersion is stronger under **bigger values of aerosol velocity** (lower values of density  $N_n$ ) → it changes threshold's dependence on the parameters, defining aerosol velocity.

For monodispersed ensemble, threshold decreases with increasing of the aerosol velocity (decreasing of  $N_n$ , increasing of  $\rho_a$ ); with the account of the dispersion, **threshold charge has a minimum over velocity**. The threshold increases with the ion diffusion coefficient.



# Mars's atmosphere (heights 70-100 km):



Neutrals:  $\text{CO}_2$ ,  $T \sim 130 \text{ K}$ ,  
 Dust particles:  $a = 1 \mu\text{m}$ ,  
 $g = 3.71 \text{ m/s}^2$

$$f_R = \frac{\exp[-(x-1)^2/\Delta^2]}{\Delta\sqrt{\pi}}$$

$x = R/a$        $\Delta = 0.05$

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$$\rho_a = 1 \text{ g/cm}^3, \delta = n_i/n_e = 10, \nu_i/N_n = 9 \cdot 10^{-10} \text{ cm}^3/\text{c}$$

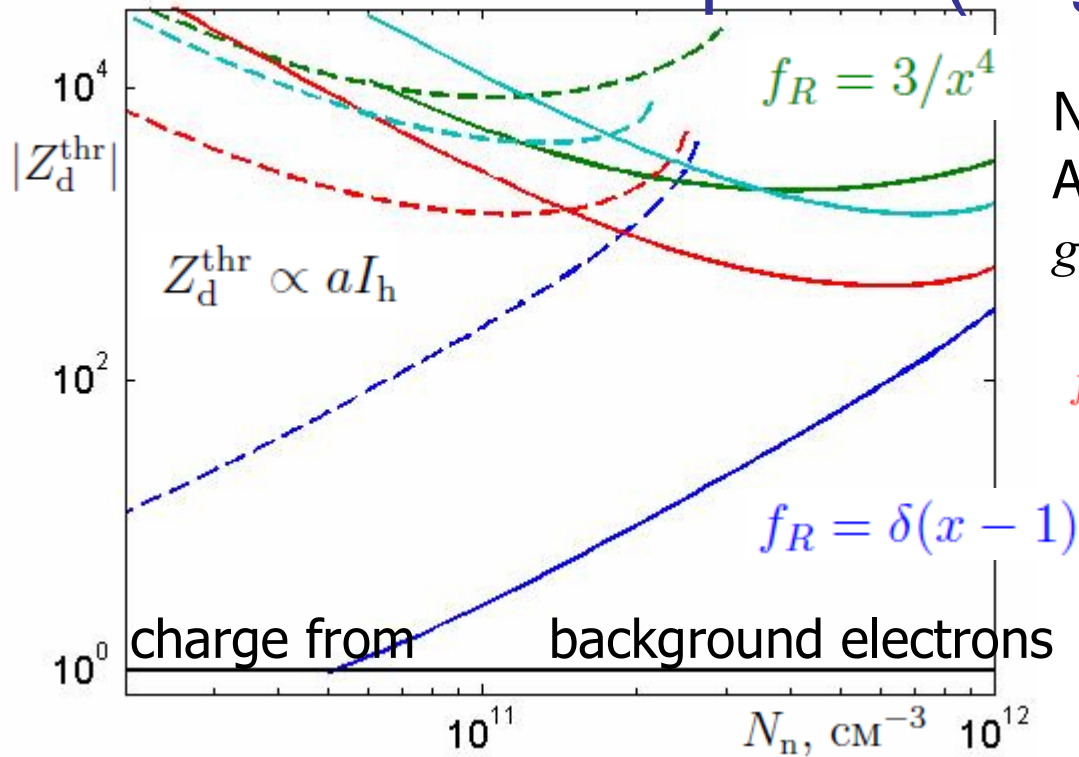

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$$\rho_a = 2 \text{ g/cm}^3, \delta = n_i/n_e = 50, \nu_i/N_n = 27 \cdot 10^{-10} \text{ cm}^3/\text{c}$$

Optimal scales:

$$\lambda_{\text{opt}} = 30 \div 150 \text{ cm} \quad \lambda_{\text{opt}} = 10^3 \div 10^4 \text{ cm} \quad \lambda_{\text{opt}} = 50 \div 1000 \text{ cm}$$

# Titan's ionosphere (heights 900-1200 km):



Neutrals:  $N_2$ ,  $T \sim 170$  K;  
 Aerosols: ice  $HC_nN$ ,  $a = 50$  nm  
 $g = 1.35$  m/c<sup>2</sup>

$$f_R = \frac{\exp[-(x-1)^2/\Delta^2]}{\Delta\sqrt{\pi}} \quad \Delta = 0.09$$

$$x = R/a \quad \Delta = 0.2$$

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 $\rho_a = 0.7$  g/cm<sup>3</sup>,  $\nu_i/N_n = 7 \cdot 10^{-10}$  cm<sup>3</sup>/c

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 $\rho_a = 1.5$  g/cm<sup>3</sup>,  $\nu_i/N_n = 21 \cdot 10^{-10}$  cm<sup>3</sup>/c

Optimal scales:

$$\lambda_{opt} = 400 \div 2200 \text{ cm} \quad \lambda_{opt} = (0.3 \div 3) \cdot 10^4 \text{ cm}$$

$$\lambda_{opt} = (0.3 \div 10) \cdot 10^4 \text{ cm} \quad \lambda_{opt} = (0.5 \div 10) \cdot 10^4 \text{ cm}$$

## Summary:

1. Size dispersion increases the instability threshold and optimal scales significantly.
2. If size dispersion has Gaussian distribution, there is no instability under big dispersions.
3. Size dispersion changes threshold's dependence on some of the parameters (threshold charge has a minimum over parameters, defining respective velocity).
4. Strict requirements for the instability to develop: small size's dispersion and a flux of high-energy electrons.

### Quantitative estimates:

	Earth	Mars	Titan
Dispersion	$( Z_d^{\text{thr}} )_{\text{min}} \approx 100 \div 450$	$( Z_d^{\text{thr}} )_{\text{min}} \approx (7 \div 30) \cdot 10^4$	$( Z_d^{\text{thr}} )_{\text{min}} \approx (450 \div 1300)$
Monodispersed ensemble	$( Z_d^{\text{thr}} )_{\text{min}} \approx 5$	$( Z_d^{\text{thr}} )_{\text{min}} \approx 8$	$( Z_d^{\text{thr}} )_{\text{min}} \approx 1$